



Contribution to the Symposium: 'Effects of Climate Change on the World's Oceans' Food for Thought

Building confidence in projections of the responses of living marine resources to climate change

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The Fifth Assessment Report of the Intergovernmental Panel on Climate Change highlights that climate change and ocean acidification are challenging the sustainable management of living marine resources (LMRs). Formal and systematic treatment of uncertainty in existing LMR projections, however, is lacking. We synthesize knowledge of how to address different sources of uncertainty by drawing from climate model intercomparison efforts. We suggest an ensemble of available models and projections, informed by observations, as a starting point to quantify uncertainties. Such an ensemble must be paired with analysis of the dominant uncertainties over different spatial scales, time horizons, and metrics. We use two examples: (i) global and regional projections of Sea Surface Temperature and (ii) projection of changes in potential catch of sablefish (*Anoplopoma fimbria*) in the 21st century, to illustrate this ensemble model approach to explore different types of uncertainties. Further effort should prioritize understanding dominant, undersampled dimensions of uncertainty, as well as the strategic collection of observations to quantify, and ultimately reduce, uncertainties. Our proposed framework will improve our understanding of future changes in LMR and the resulting risk of impacts to ecosystems and the societies under changing ocean conditions.

Keywords: climate change, fisheries, marine resources, multi-model ensembles, projection, uncertainty.

Living marine resources projections under climate change

The Fifth Assessment Report of the Intergovernmental Panel on Climate Change (IPCC AR5) highlights that changes in ocean temperature, oxygen, carbonate system, and other ocean properties are contributing to the challenges of sustainable ocean management (Field *et al.*, 2014). The importance of a comprehensive assessment of the impact of climate change on the ocean is highlighted by two

new ocean-specific chapters within the IPCC AR5 Working Group II (WGII) on impacts, adaptation, and vulnerability (Field *et al.*, 2014). In relation to living marine resources (LMRs), the IPCC Report concludes with medium to high confidence that marine species have been shifting their ranges, seasonal activities and periodicities, migration patterns, abundances, and inter-/intra-specific interactions that result in changes in trophodynamics in response to changing ocean conditions (Pörtner *et al.*, 2014). These changes

are projected to lead to altered patterns of ocean productivity, biodiversity, and fisheries catch potential in the 21st century (Kirby and Beaugrand, 2009).

One of the advances in assessing the impacts of climate change on LMR in the IPCC AR5 WGII over previous assessment reports is the wider availability and use of ecosystem model projections. These quantitative model projections include shifts in net primary productivity, the distribution of exploited populations and changes in potential fisheries production and ecosystem structure at local and global scales (Pörtner et al., 2014). Projections have been generated from modelling approaches that range from global coupled atmosphere-ocean-biogeochemistry earth system models (e.g. Bopp et al., 2013), to species distribution models (e.g. Cheung et al., 2009), single-species population dynamic models (e.g. Lehodey et al., 2010), and whole ecosystem models (e.g. Ainsworth et al., 2011; Griffith et al., 2011). The scope, objectives, assumptions, scales (spatial and temporal), and degree of validation with empirical data vary widely across these models, and approaches range from highly empirical to highly mechanistic (Barange et al., 2010; Fulton, 2010; Plagányi et al., 2011; Stock et al., 2011).

Statements of confidence concerning the impacts of climate change on LMRs within the IPCC AR5 WGII report were based on a qualitative assessment of observational evidence and individually published projections encompassing the diversity of LMR models described above. While this is a necessary starting point, more quantitative confidence estimates for projections can increase their utility for policy formulation and evaluation. There is therefore a need for a quantitative framework for systematically exploring uncertainties in LMR projections. Such a framework would also help identify where investment in further theoretical development, observational measurements, and model development are needed, ultimately improving the reliability of climate-LMR projections (Cheung et al., 2013a; Brander, 2015). Systematic exploration of uncertainties have been undertaken for climate and oceanographic projections (e.g. the Atmospheric Model Intercomparison Project (Gates, 1992) and the Coupled Model Intercomparison Project (Meehl et al., 2000; Taylor et al., 2011) and for impact assessments of selected sectors (e.g. Agricultural Model Intercomparison and Improvement Project; Rosenzweig et al., 2013). Exploration of uncertainties is also an important component in traditional fisheries resource assessment, while the increasing demand for ecosystem-based fisheries management raises additional challenges to systematically understanding projection uncertainties (e.g. Hill et al., 2007; Link et al., 2012). More recently, initiatives on comparing fisheries models (e.g. Fisheries Model Intercomparison Project, ICES-PICES Strategic Initiative on Climate Change Impacts on Marine Ecosystems) have also been started.

While challenges in quantifying uncertainty in climate-LMR projections for global change assessment parallel those considered in modelling other complex natural systems such as climate, there are additional sets of complexity that are specific to LMRs. Climate-LMR projections require linking physical, biological and human sub-systems across different temporal and spatial scales. Such inter-linkages lead to additional uncertainties that originate in particular systems or scales (Planque, 2016). In addition, the behaviour of some components of LMR systems is difficult to predict, such as the responses of fishing activities to changes in climate and fisheries resources. Moreover, many LMR models require large number of input parameters relative to the available observational data that are available to calibrate and validate the model outputs. Techniques for assessing model uncertainties that are commonly

applied to conventional fisheries assessment (e.g. Bayesian estimates of process and observation errors) are thus difficult to apply to climate-LMR projections.

This paper aims to synthesize our knowledge of the uncertainties of LMR projections under climate change and propose a framework to systematically assess such uncertainty. Our paper complements that of Payne et al. (2016), which reviews existing approaches in addressing uncertainties in LMR. Here, we focus on the following: first, we characterize different types of uncertainty in climate and LMR projections, highlighting the challenges of the large uncertainty space; second, drawing from the experience of physical climate model intercomparisons, we explore how multi-model comparison and ensemble frameworks can be used to systematically identify and quantify uncertainties in LMR projections. Through an example, we highlight the relative roles of uncertainty linked to climate variability, climate model uncertainty, and future emissions scenarios as a function of time horizon and spatial scale. This is followed by a discussion of the role of observations in refining uncertainty estimates. Finally, we discuss how outcomes from this model-assessment framework can be used to evaluate the risk of climate change to LMRs and inform the design of management and conservation measures to reduce such risk.

Sources of uncertainty

Climate-LMR models that estimate the impacts of climate change generally have three model components that are linked to describe the responses of marine resources, fisheries, and human society to climate systems. These components generally include an atmosphere-ocean-biogeochemical and lower-trophic level models, a fish or upper-trophic level model (Holt et al., 2014), and a model for the extraction and availability of ecosystem services from marine ecosystems (see Fulton, 2010; Plagányi et al., 2011; Stock et al., 2011). The three components are either related “off-line”, where each model component is run separately with the outputs from one component used as inputs for another (Cheung et al., 2011; Blanchard et al., 2012; Christensen et al., 2015), or dynamically (i.e. “on-line”) with the models incorporating fully interactive processes and, in some cases, feedbacks among the three components (Fulton, 2010; Lefort et al., 2015).

Research on physical climate projections, biodiversity, and ecological modelling has recognized many topologies of uncertainties (Regan et al., 2002; Link et al., 2012). Modelling of physical and biogeochemical properties of atmospheric and ocean systems in climate change assessments have commonly categorized uncertainties, for any time horizon and spatial scale, into three components: (i) internal variability, (ii) model uncertainty, and (iii) scenario uncertainty (Table 1) (Hawkins and Sutton, 2009). In our discussion of the uncertainties associated with climate-LMR projections, we adopt this terminology to leverage the knowledge and experience of the climate modelling communities.

Internal variability is caused by natural physical and ecological processes that are intrinsic to climate and ecological systems. It arises in both temporal and spatial dimensions, even in the absence of any external (e.g. anthropogenic) perturbations, and includes phenomena such as the *El Niño*-Southern Oscillation (ENSO), the North Atlantic Oscillation (NAO), the Atlantic Multidecadal Oscillation (AMO), variations in gyre boundaries not correlated with major climate models, and predator-prey cycles, etc. (Day, 1982). Century-scale climate change projections developed in association with the IPCC realistically resolve many modes of internal climate variability, but these simulations are not designed to simulate

Table 1. Summary of different types of uncertainties in LMR models.

Types of uncertainties	Description	Examples	
		Coupled atmospheric, ocean and biogeochemical models	Fish and fisheries models
Internal variability	Natural variations of physical, biogeochemical, and ecological processes that contribute random variability to projections of LMRs	ENSO and PDO, NAO, AMO, locations of gyre boundaries	Predator–prey dynamics, spatial and temporal variations in fish populations not arising from deterministically modelled climate change signal
Model uncertainty			
(a) Parameter	Specific parameter values used in the formulae determining the behaviour of the models	Parameters controlling sub-grid scale oceanographic processes, phytoplankton growth, zooplankton grazing, biogeochemical transformations, and detritus remineralization	Values of the parameter describing diet composition, dispersal rate, production and consumption rates, trophic interactions, and other ecological/anthropogenic processes represented in the models. If variations in parameter values reflect an alteration of model architecture or design, it should belong to the structural uncertainty category
(b) Structure	Differences in abstraction, understanding and representation of the system through different model architecture, design and assumptions, the method of representing space/time, and the kinds of ecological processes, human and natural drivers included	Grid resolution; the number of nutrients included in biogeochemical models, and differences in representing foodweb structure	Size-based vs. functional group-based approaches; the number of functional groups included; differences in representing foodweb structure, fish movement and different life history stages in fish models
Scenario uncertainty	Differences in the natural and/or anthropogenic forcing that drive the model simulation	Representative Concentration Pathways developed by the IPCC	Shared socio-economic pathways, spatial and temporal changes in fisheries, and how all these influence model components directly or indirectly

a specific observed event or predict a future event, and will not capture all aspects of spatial and temporal scales of these modes (Guilyardi *et al.*, 2009). For ecological systems, natural fluctuations that are driven by environmental variability and dynamics of ecological interactions are often difficult to predict (Beckage *et al.*, 2011; Deser *et al.*, 2012), causing systematic or seemingly random variations in ecological states that may persist for a decade or more (Deser, 2012; Stocker *et al.*, 2014). Different initial conditions of the climate or LMR models, representing different realizations of the climatic and ecological systems, will generate different patterns of internal variability. Thus, one method to explore internal variability is to analyse simulation results generated from ensemble members of climate and ecological models that have different initial conditions.

Model uncertainty is made up of two sub-categories: parameter and structural uncertainty (Tebaldi and Knutti, 2007). Parameter uncertainty relates to the specific parameter values used in the formulae that influence the behaviour of a model (Tebaldi and Knutti, 2007; Knutti *et al.*, 2010). For parameters that are estimated from observations, parameter uncertainty stems from our limited ability to precisely measure or estimate specific physical or ecological processes and quantities (Link *et al.*, 2012), as well as from the inherent variability in certain processes (e.g. growth rates that vary across individuals) that are not resolved within the models.

Structural uncertainty relates to the spatial, temporal, and mathematical resolution employed by a model and the types of processes that are represented. Structural uncertainty includes the function forms of equations used to describe mechanistic processes and the types of interactions assumed to influence climate-LMR processes. Such uncertainties cannot be explored via parameter perturbations. For example, explicit trophic relationships that are not described by size-structured interactions are not represented in size-based trophodynamic models (e.g. Blanchard *et al.*, 2012; Watson *et al.*, 2014), while such relationships may be included in functional group type foodweb models (e.g. Christensen and Walters, 2004).

Scenario uncertainty relates to the many possible futures comprising different socio-economic policies and technological developments likely occurring over the course of a model projection (e.g. Moss *et al.*, 2010; Nakicenovic *et al.*, 2014). Climate-LMR model drivers include the spatial and temporal changes in greenhouse gas and aerosol concentrations, fishing effort, and other human social-economic activities. Scenario uncertainty is not completely independent of internal variability in the climate-LMR system, as future decisions on the utilization and conservation of resources are sensitive to natural variation in the availability and distribution of LMR (e.g. the fishing quota decided on for the next management cycle are dependent on the productivity and abundance of the resources, as well as on how neighbouring countries or regions are managing their resources).

The full range of possible future states for a given LMR reflects contributions from all the sources of uncertainty outlined above, with potential cascades of uncertainties interacting and accumulating over components of the climate-LMR models (Figure 1). For any particular scenario, LMR models that differ in their structure and parameter values will simulate a range of future changes in ocean biogeochemistry, fish and fisheries. Additionally, an individual model with a fixed set of parameters will display variability in projections as a result of the internal variability associated with natural fluctuations of the climatic or ecological systems. Uncertainties that originate in different climate-LMR model sub-components may be additive or multiplicative. Thus, the final scope of uncertainties of

LMR projections is expected to be different from the uncertainty scope of each model sub-component.

The width of the envelope of uncertainty depends on the nature of interactions between linked models; the types of interactions include linearity of the linkages, existence of threshold responses, and positive/negative feedbacks (Peters and Herrick, 2004). When the processes linking two or more models are non-linear, uncertainties may be dampened or magnified through model linkages, for example, through attenuation or amplification of changes in higher trophic level production in marine ecosystems driven by climate change (Chust *et al.*, 2014; Stock *et al.*, 2014a). Feedbacks in social-ecological systems can be positive or negative, and uncertainties propagated in models that are linked dynamically with feedbacks resulting in emergent dynamics are difficult to predict.

Here, we draw experience from the large body of research on exploring uncertainties of climate projections to propose that the envelope of uncertainties of climate-LMR projection can be explored by systematically quantifying the three categories of uncertainty that we discussed above: internal variability, model uncertainty, and scenario uncertainty. Review on specific techniques to explore each source of uncertainties can be found in Payne *et al.* (2016).

Experiences from quantifying uncertainty of climate projections

For ocean-atmospheric general circulation models and biogeochemical models, the Coupled Model Intercomparison Project Phase 5 (CMIP5) multi-model database allows assessment of uncertainty in climate change projections across the dimensions illustrated in Figure 1. Climate change projections were produced from >30 models, developed by different modelling groups with a standard set of scenario experiments (Flato *et al.*, 2013). The CMIP5 database allows some exploration of uncertainty, but comprehensive categorization of uncertainty into structural uncertainty, parameter uncertainty, and internal variability is not possible. The main challenges include the limited number of modelling groups that were able to contribute ensembles of runs, some models are fully independent of each other, and a lack of exploration of parameter uncertainty. Ideally, the ensemble should consist of a random sample across the uncertainty components in Figure 1. For complex inter-linked models such as climate or climate-LMR models, exploring their full scope of uncertainty would require substantial computational time and other resources. Thus, a systematic approach is needed to efficiently explore the envelope of uncertainties.

To further explore the uncertainty contributed by internal variability for each model, ensembles of climate simulations have been run under identical forcing, but with each simulation initialized with slightly different, but equally plausible, conditions (Rodgers *et al.*, 2015). The chaotic nature of climate variability quickly produces different climate trajectories in each ensemble member (Wittenberg *et al.*, 2014). By considering each of the trajectories as a plausible outcome, the ensemble can be used to isolate that part of projection uncertainty due to internal variability (Frölicher *et al.*, 2009; Deser *et al.*, 2012).

Hawkins and Sutton (2009) analyse CMIP3 (i.e. the precursor of CMIP5) projections to explore the contribution of internal variability and model and scenario uncertainties to climate projections at global and regional scales. They showed that the dominant sources of uncertainty in surface air temperature projections vary with spatial scale and time horizon, noting the importance of model uncertainty and

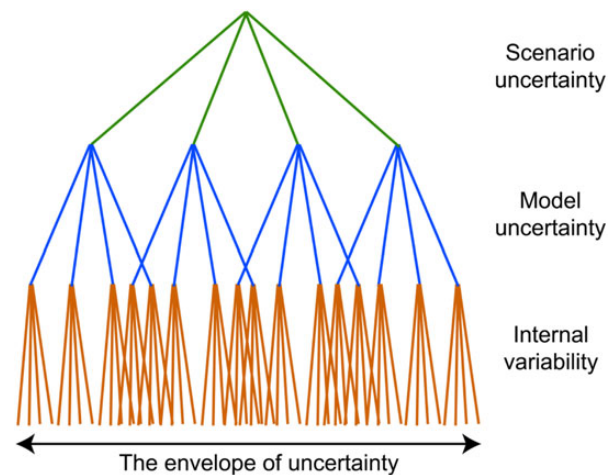


Figure 1. Schematic example illustrating cascades of uncertainties in LMR projection (modified from Wilby and Dessai, 2010). For a particular time horizon and spatial scale, the range, or envelope, of possible outcomes includes contributions from scenario uncertainty (green), model uncertainty (blue), and internal variability (orange). The cascades of uncertainties apply to each of the sub-components of climate-LMR models. Uncertainties from each model sub-component may be additive or multiplicative. In this schematic diagram, the width of each uncertainty level does not imply the magnitude of the uncertainty. For example internal variability may be larger than scenario uncertainty and vice versa. This figure is available in black and white in print and in colour at ICES *Journal of Marine Science* online.

internal variation for mid-21st century regional projections. To further illustrate the application of the framework used by Hawkins and Sutton (2009) in the oceanic realm, we analysed the projection uncertainties for sea surface temperature by combining CMIP5 projections and a large ensemble projections from the Earth System Model of the Geophysical Fluid Dynamic Laboratory (GFDL ESM2M model; Dunne *et al.*, 2012, 2013; Rodgers *et al.*, 2015).

We used the projection of SST as an example of exploring the sensitivity of model projections to different sources of uncertainties. Scenario uncertainty is estimated to be the difference between the multi-model mean of projections from 15 CMIP5 models of two 21st century emissions scenarios: the low-emissions scenario RCP2.6 with an increased radiative forcing that peaks at $\sim 3 \text{ W m}^{-2}$ before 2100 then declines to 2.6 W m^{-2} by 2100, and the high-emissions scenario RCP8.5, with an increased radiative forcing of $> 8.5 \text{ W m}^{-2}$ by year 2100 (Meinshausen *et al.*, 2011). Model uncertainty is estimated as the standard deviation of changes in SST (10-year running mean) from each model projections. The internal variability is estimated as the standard deviation of projections from 30 ensemble member simulations of GFDL ESM2M (Rodgers *et al.*, 2015).

Globally, the analysis shows that model uncertainty is dominant in the medium term SST projection (2030–2050), while the long-term (2080–2100) projection is dominated by scenario uncertainty (Figure 2). The large model uncertainty over the medium term reflects the large variations in regional scale biases in the models. Although the importance of internal variability is second to model uncertainty in near term projection (2010–2030), its relative importance decreases rapidly further into the future.

The relative importance of different uncertainty sources also varies between different regions. In the Northeast Atlantic (North Sea Large Marine Ecosystem; Pauly *et al.*, 2008), the importance of scenario uncertainty is smaller compared with those projections

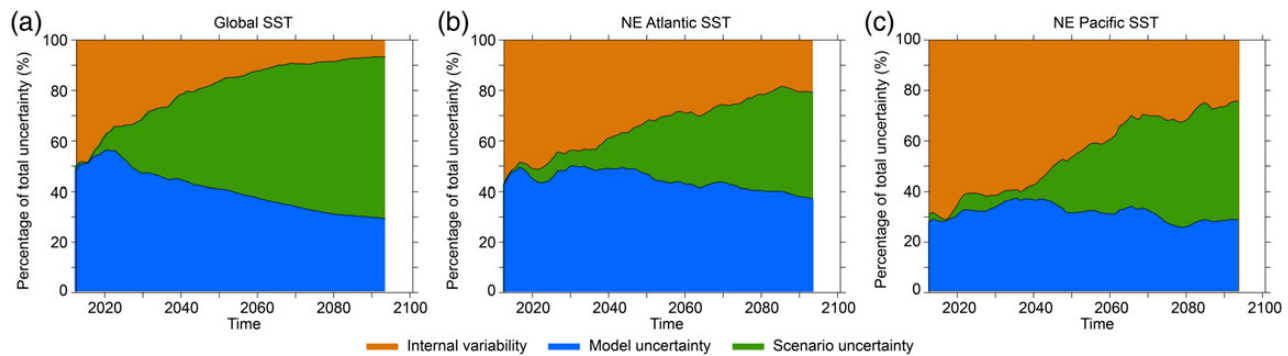


Figure 2. The relative importance of each source of uncertainty in annual mean sea surface temperature projection is shown by the fractional uncertainty for (a) global mean, (b) Northeast Atlantic, and (c) Northeast Pacific in the 21st century. Uncertainties are separated into three components: internal variability (orange), model uncertainty (blue), and scenario uncertainty (green). The percentage of total uncertainty is calculated from dividing the level of uncertainty from the specific component by the sum of the three types of uncertainties. For internal variability, the standard deviation of annual mean SST from the GFDL ESM2M ensemble is calculated year by year. The same procedure has been applied for model uncertainty, but a 10-year running mean (longer than the typical ENSO period) is first applied to the individual CMIP5 model projections. This figure is available in black and white in print and in colour at *ICES Journal of Marine Science* online.

at the global scale, while model uncertainties and internal variability become the dominant uncertainty sources. The internal variability in the Northeast Atlantic may represent known properties of interannual and multidecadal climate and oceanographic variability such as NAO and AMO (Viles and Goudie, 2003; Beaugrand and Kirby, 2010). In the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem), internal variability becomes a dominant source of uncertainty representing properties, such as ENSO and Pacific Decadal Oscillation (PDO). In both basin-scale examples, the internal variability of SST is a bigger contribution to projection uncertainty than in the global scale projection (Figure 3). Moreover, in the short to medium term, the projected increase in SST is not sensitive to different emission scenarios, both globally and in the NE Atlantic and NE Pacific (Figure 3). However, long-term warming is much more sensitive to different emission scenarios, particularly at the global scale. We anticipate the increased importance of internal variability observed at the basin scale may be even more prominent when examining smaller spatial scales.

In addition to highlighting the relative contribution of different sources of uncertainty, this exploration of uncertainty suggests strategies to prioritize investment to improve our understanding of specific types of uncertainties. In the example presented here, the large model uncertainties in the projections of SST in the Northeast Atlantic call for better understanding of key processes that may be represented differently among models. In the Northeast Pacific, where large internal variability is difficult to reliably predict, the medium term effects of greenhouse gas emission will be difficult to separate from natural variability. This further highlights the need for better understanding interannual variability and thus the need for longer term observational records.

Systematic exploration of climate-LMR projection uncertainties

Systematic exploration of the components of uncertainty in both space and time dimensions in a manner analogous to examples from physical climate model projections (Figure 3) is critical for moving quickly toward refined uncertainty bounds on climate-LMR projections. Thus, exploration of uncertainties within climate-LMR projections would include: (i) making projections from ensemble members of

models with different properties of internal temporal or spatial variability, (ii) making projections from ensemble members of models with different model structure and parameter values, and (iii) generating projections that are based on different climate and fishing scenarios.

The conditions to systematically explore uncertainties within climate-LMR projections already exist. For fish and fisheries models, attempts to explore the full matrix of uncertainties (particularly model uncertainty with scenario uncertainty) have been made for a limited number of fisheries or stocks (Table 2). Existing examples mainly involve Management System Evaluations in which the performance of different models is assessed under different management scenarios (Link *et al.*, 2012). Methods such as Monte-Carlo simulation, Bayesian statistical frameworks, and a plethora of quantitative methods also provide a basis for exploring both the parameter and structural components of model uncertainty (Hill *et al.*, 2007; Hollowed *et al.*, 2013). Moreover, various statistical approaches are available to analyse the properties of different components of uncertainty, and how they contribute to the full scope of uncertainty (Saltelli *et al.*, 2000). Furthermore, initiatives such as the fisheries component of the Inter-Sectoral Impact Model Intercomparison Project (Warszawski *et al.*, 2014), which aim to develop LMR projection databases for climate-fisheries assessment that are similar in nature to CMIP now been established. Such a database would facilitate collaborative efforts of LMR research communities to explore the full scope of uncertainties.

A remaining knowledge gap in climate-LMR uncertainty exploration is the limited understanding of uncertainties arising from internal variations in the ecological system or fishing scenarios in projecting LMR changes, as well as their interactions with internal variability at different temporal and spatial scales. The linkages between physical and biogeochemical ocean changes and ecosystem responses are likely to be non-linear and may also involve thresholds; thus the resulting pattern of internal variability of climate-LMR model projections are likely to be more complex. For example, the actual response of LMRs to a particular level of environmental change may be limited by predator–prey interactions, or altered by species-specific sensitivity and adaptability to environmental fluctuations (Foden *et al.*, 2013).

Exploration of internal variability in climate-LMR projections can be done by comparing projections from ensemble members of

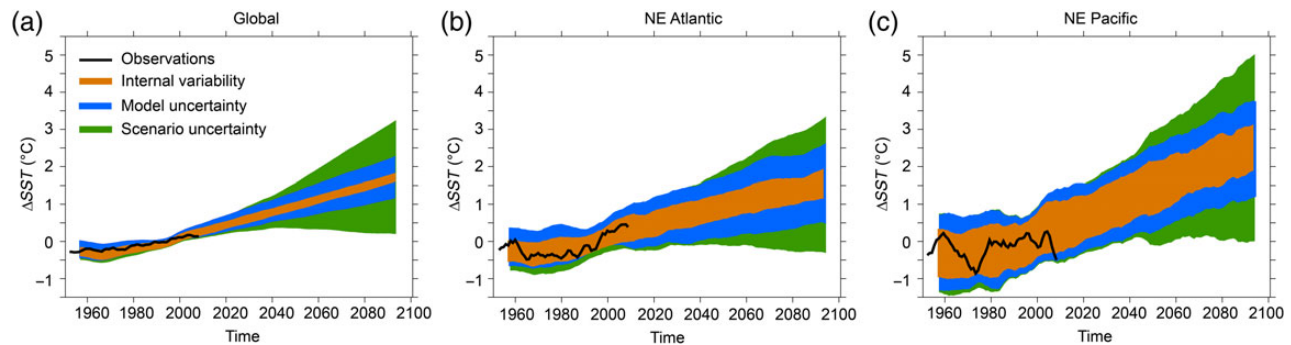


Figure 3. Changes in annual average sea surface temperature (10-year running mean) for (a) global mean, (b) Northeast Atlantic, and (c) Northeast Pacific relative to the 1986–2005 mean. SST observations (black line) are based on Smith *et al.* (2008). The uncertainty area was calculated by adding and subtracting the errors from each uncertainty source (internal variability, model uncertainty, scenario uncertainty) to and from the ensemble-mean projection of 15 CMIP5 models. Errors from different uncertainty sources are assumed to be additive. This figure is available in black and white in print and in colour at ICES *Journal of Marine Science* online.

a single model with different sets of initial conditions. For example, we used three versions of Dynamic Bioclimate Envelope Model (DBEM) (Cheung *et al.*, 2011) to project changes in maximum potential catch of sablefish (*Anoplopoma fimbria*) in the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem) from 2000 to 2060 (Figure 4). Specifically, we explored the effects of internal variability of ocean conditions using 20 different ensemble member projections from the GFDL ESM2M (Rodgers *et al.*, 2015). We also compared the relative contribution of uncertainties from internal variability of ocean conditions, structural uncertainties of DBEM, and uncertainty from different climate scenarios (RCP 2.6 and 8.5).

The results suggest that internal variability is a dominant source of uncertainty for sablefish in the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem) by 2060 relative to 2000, followed by the structural uncertainties of DBEM. Scenario uncertainty contributes <10% of the total uncertainty. This is broadly consistent with the projected SST changes in this region, with internal variability contributing ~40–70% of the total uncertainty over the time frame of 2000–2060 (Figure 2). However, model uncertainty is substantially lower for sablefish projections relative to SST projections, possibly because the structural difference between CMIP5 models (used in SST projection) is much larger than those between the three versions of DBEM (used in sablefish projection). Also, in addition to SST, DBEM projections are driven by other ocean-biogeochemical variables, such as oxygen and net primary production (Cheung *et al.*, 2011). Internal variability of multiple oceanographic properties may have magnified the internal variability of the DBEM projections.

Since DBEM outputs represent mainly long-term trend of potential catches, interannual variation of reported catches is substantially higher than the internal-variability of the projections (Figure 5). DBEM does not represent some processes that contribute to interannual variability of catches such as recruitment variability and changes in fishing effort. Besides spawning stock abundance, recruitment variability could be dependent on both physical (temperature, wind, and current) and/or biological (primary productivity, predation pressure) at different spatial and temporal scales (Houde, 2008). The relative importance of these factors and the processes contributing to recruitment vary between species. In addition, catches are also dependent on changes in fishing effort which can be dependent fisheries management (e.g. quota), social-economics factors (e.g. price of fish and cost of fishing),

and fishers' behaviour. DBEM does not resolve many of these processes and does not have species-specific recruitment submodel. Therefore, DBEM is not expected to represent the actual interannual variability of the catch. On the other hand, DBEM is structured to represent the long-term trends of resource productivity. The long-term trend (20-year running mean) of the reported catch of sablefish falls within the range of trajectories of the projections (Figure 5).

The example of the sablefish highlights the need to carefully consider the actual processes that are represented by the sample of LMR models in quantifying uncertainty from model ensembles. This challenge applies to both ocean biogeochemical and LMR models (e.g. Jones and Cheung, 2015; Cabré *et al.*, 2015). For instance, the relatively coarse-resolution Earth System Models do not capture potentially large random variability associated with submesoscale and mesoscale ocean features such as fronts, eddies, and filaments (Stock *et al.*, 2011).

A standardized set of climate-LMR scenarios is needed to quantify scenario uncertainty for climate-LMR projections. These scenarios must be reconciled with a range of different realizations of future emission (e.g. IPCC AR5's Representative Concentration Pathways, RCPs) (Moss *et al.*, 2010) and social-economic development (e.g. Shared Socio-economic Pathways or the Sustainable Development Goals) (Griggs *et al.*, 2013; Hunter and O'Neill, 2014). However, emissions scenarios only describe broad-brush societal changes in the 21st century. Fishing sector-specific storylines concerning management, aquaculture, and technological development, and demand for fish in countries across the economic development spectrum at global and regional scales are also needed. Such factors would ultimately affect the magnitude and distribution of fishing effort. Trajectories of other human marine-related activities that drive changes in marine ecosystems should also be included (Figure 6). Development of these scenarios requires interdisciplinary collaboration between natural and social scientists (Österblom *et al.*, 2013). Although, there are currently independent efforts to develop such scenarios at global and regional scales (e.g. Barange *et al.*, 2014; Jones *et al.*, 2015), community-wide effort in developing standardized sets of scenarios would facilitate consistent comparison of LMR projections.

Building confidence and constraining the scope of plausible projections with observations

Observations across different scales are critical for building confidence in projections and reducing the scope of LMR uncertainty

Table 2. Selected case studies that explored different aspects of uncertainties in projections of aquatic (marine and freshwater) biological resources under climate change.

Spatial scale	Selected case studies	Explored uncertainties and conclusions
Global	Variability of projections of distribution and patterns of species turnover across three different species distribution models for over 800 commercially exploited fish and invertebrates in the world under two greenhouse gas emission scenarios (Jones and Cheung, 2015).	<ul style="list-style-type: none"> – Structural uncertainties of species distribution models – Scenario uncertainties of greenhouse gas emission pathways – Larger variability in projections exists between greenhouse gas emission scenarios (RCP2.6 and RCP8.5) than between three different species distribution models
Regional (UK waters)	Projecting changes in maximum catch potential and profitability from fishing 31 key commercially targeted fish species primarily inhabiting UK waters using different climate models, species distribution modelling approaches, and socio-economic scenarios (Jones et al., 2015). Three fisheries and socio-economic scenarios were designed based on key variables identified in the Alternative Future Scenario for Marine Ecosystems scenarios	<ul style="list-style-type: none"> – Structural uncertainties of species distribution models and climate models – Scenario uncertainties of greenhouse gas emission, fisheries, and socio-economics pathways – Scenario (climate, fisheries, and socio-economic) uncertainty dominates over structural uncertainty of climate and biological models
Regional (Central North Pacific Ocean)	Uncertainty of a trophodynamic model (Ecopath with Ecosim) was explored using Monte-Carlo simulation. Confidence limits of key input parameters were set based on the reliability of the data, as indicated by the data type. Results from 500 dynamic simulations (each involving up to several thousand iterations to find a balanced model) were used to construct 95% confidence intervals for the derived biomass time series (Kearney et al., 2012)	<ul style="list-style-type: none"> – Parameter uncertainty of the ecological models
Regional (Eastern US coast)	Using experimentally derived thermal tolerance limits to project range shift of grey snapper (<i>Lutjanus griseus</i>) in estuaries along eastern US coast. Projections were driven by temperature simulated from 23 different climate models, two thermal tolerance metrics under three different emission scenarios (Hare et al., 2012)	<ul style="list-style-type: none"> – Parameter uncertainty of range shift model – Structural uncertainties of climate models – Scenario uncertainties of greenhouse gas emission pathways – Different species distribution models contributed the largest variation in projections, followed by different GCMs. The contribution of variability from different GCMs increased over time and to a level that is comparable with variability from different species distribution models for end of 21st century projections. Different observation datasets had a small influence on the overall variability of the projections
Regional (freshwater ecosystems in France)	Projection of distribution shifts of 35 species of freshwater fish in France across 100 random subsets of observation data, seven species distribution models, and climate projections from 12 climate models, resulting in 8400 different potential futures projections (Buisson et al., 2010)	<ul style="list-style-type: none"> – Parameter uncertainty of species distribution models – Structural uncertainties of species distribution models and climate models – Scenario uncertainty of greenhouse gas emission pathways – Uncertainty about thermal limits of the species dominates over model or scenario uncertainties

by constraining parameters, model structures, and eliminating implausible solutions. Model metrics derived from observations can be compared with model outputs to obtain a quantitative assessment of model skill. Model metrics that are of particular interest for LMR models include species distributions, community structure, abundance and fisheries catches (Table 3). These data are generally available for broad-scale evaluation.

Different LMR models may vary in their ability to represent seasonal cycles, interannual, and/or long-term (decadal or longer) trends, with skill at one scale not always implying skill at others (Table 3). To assess confidence in the temporal properties of climate-LMR projections, we suggest three possible tiers of evaluation that involve the use of observational data to assess consistency with: (i) mean observed spatial patterns or seasonal climatologies across the scale of interest, (ii) previously observed responses to climate variability, and (iii) observed long-term

trends attributable to climate change, fishing, and other human drivers. In practice, the ability to assess models across all three tiers is often limited by data availability, particularly the paucity of the long-term, comprehensive, and high-quality datasets required to assess models against the often subtle trends in tier 3. Comparisons between LMR data and model projections are also challenging due to issues of consistency between time frame and spatial scales, as well as the confounding effects of multiple human pressures, such as climate and fishing (McOwen et al., 2014). These challenges should not, however, preclude improving confidence in LMR-climate projections.

Confidence in climate-LMR model projections can arise from model evaluation across a subset of tiers, as well as the reliance of models on robust physiological and ecological principles (Stock et al., 2011). Real caveats, however, are needed. Observational limitations also suggest that great care should to be taken

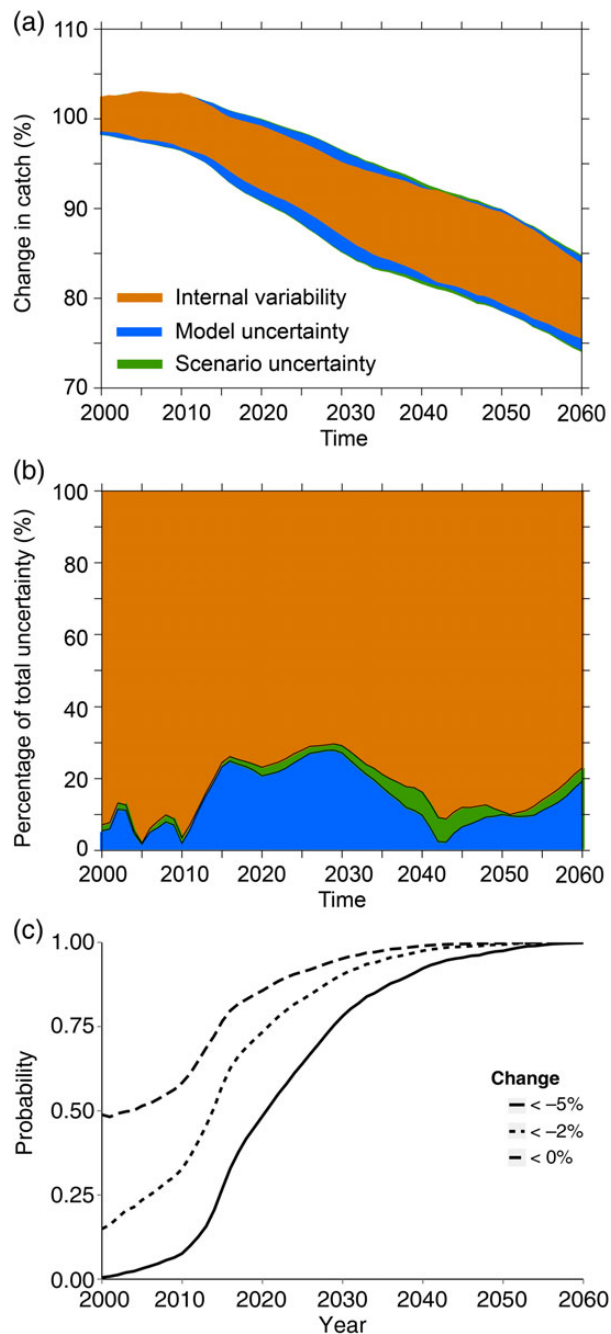


Figure 4. Projected changes in maximum potential catches of *Anoplopoma fimbria* from 2000 to 2060 under climate change. The projections were generated from using three versions of DBEM (Cheung *et al.*, under review), driven by outputs from GFDLES2M. Internal variability was estimated from projected changes in catch potential driven by outputs from 20 ensemble members of GFDL ESM2M (Rodgers *et al.*, 2015). (a) Projected changes in maximum potential catch and their standard deviation resulting from the three different types of uncertainties. (b) The relative contribution of each type of uncertainty, expressed as the proportion of total uncertainty, and (c) the probability of projecting a decrease in catch potential for >0% (dashed line), 2% (dotted line), and 5% (solid line). Model uncertainty represents variation of projections from the three versions of DBEM. Scenario uncertainty represents variations in projections between RCP2.6 and RCP8.5. This figure is available in black and white in print and in colour at ICES Journal of Marine Science online.

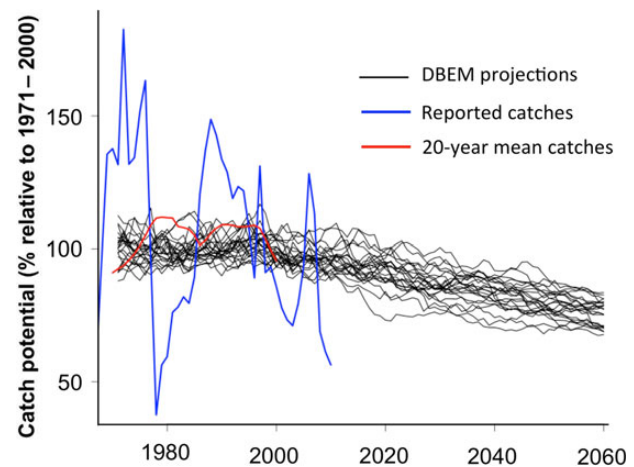


Figure 5. Comparison between projected changes in annual mean catch potential (relative to 1971–2000) using the three versions of DBEM and 20 GFDL ESM2M ensemble members under the RCP8.5 scenarios with the reported catches (from SAU: www.seararoundus.org) of sablefish in the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem). Reported catches are also smoothed by a 20-year running mean.

eliminating particular projections from consideration within an ensemble framework. Thus, a coarse culling of grossly inconsistent simulations (Overland *et al.*, 2011) is suggested rather than attempting to finely weight models based on nuanced differences in model-data fit. Even if model projections fit well with observational data, it does not guarantee that the model can accurately predict future changes, particularly when future conditions (environmental conditions or human activities) lie outside the bounds of historical conditions. In addition, a good fit between model projections and observational data could, on occasion, be more indicative of over-parameterization rather than prediction skill.

Observation data and model metrics

In the paragraphs that follow, we review available LMR data and their potential use as model metrics for evaluating LMR projections across the three tiers of evaluation described previously. We focus on the utility of three broad categories of LMR observations: fisheries-dependent data, scientific surveys, and species occurrence records. Similar efforts focusing on metrics for physical climate models (Knutti and Sedláček, 2013) and biogeochemical/plankton foodweb models (Stock *et al.*, 2014b) are also being undertaken. We also identify key uncertainties associated with such observational data, as these would complicate their use in assessing the reliability of LMR projections.

Fisheries-dependent data

Fisheries catch data are particularly useful for Tiers 1 and 2 evaluations as they are of direct relevance to LMRs and their broad spatial, temporal, and taxonomic coverage. Total catch potential can be estimated from the maximum catch of historical time-series, under certain assumptions concerning fishing effort (Cheung *et al.*, 2008; Friedland *et al.*, 2012). Moreover, spatial patterns and temporal changes in catch volume (Cheung *et al.*, 2013c) and functional and taxonomic composition (Cheung *et al.*, 2013b) of fisheries catch can be obtained from global fisheries databases. Species composition can be aggregated by body size classes (for size-based LMR models; Blanchard *et al.*, 2012), functional role (for functional

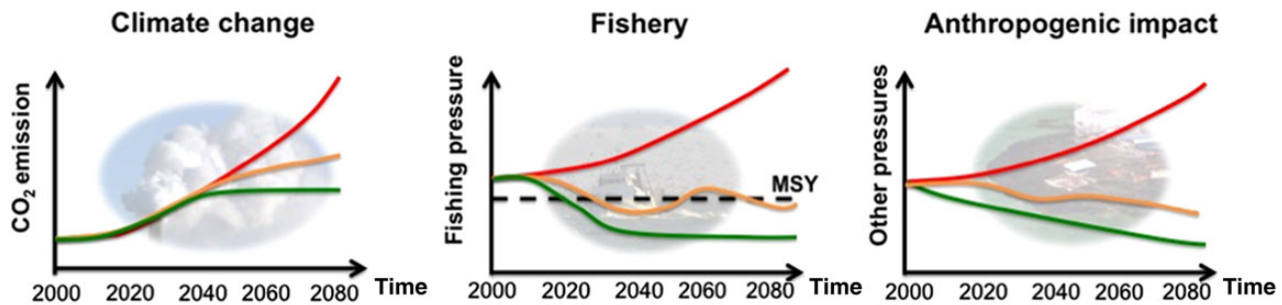


Figure 6. Schematic diagram showing an example of potential standardized sets of scenarios to be developed to explore scenario uncertainties. The red, yellow, and green lines represent different scenario pathways to be explored by climate-LMR models. Anthropogenic impacts may include contaminant level, invasive species, and habitat change. This figure is available in black and white in print and in colour at *ICES Journal of Marine Science* online.

group trophodynamic models; Christensen and Walters, 2004), and by species (for species distribution models; Cheung *et al.*, 2011). Fisheries catch data can be obtained from the Sea Around Us (SAU) project (www.seaaroundus.org), which provides spatially explicit estimates of global catches from 1950 onward. In addition, the recent effort of SAU to reconstruct catches that are not reported in the United Nations Food and Agriculture Organization landings statistics further improves the utility of such data for use as a metric for model comparisons (e.g. Zeller *et al.*, 2006). For example, in the Northeast Atlantic, fisheries catch and effort data since the early 20th century can be used to understand the ability of LMR models to reproduce changes driven by the Atlantic Multidecadal Oscillation and the NAO (Kerby *et al.*, 2013). Similar examples of the potential use of long-term series of fish and fisheries data are also available in the Northeast Pacific (Lindgren *et al.*, 2013), and large pelagic longline catch data are also available for ocean basins. As such datasets are spatially explicit, the estimated catch per unit effort can be used as an indicator of the distribution of large pelagic fish, including tunas, billfish, and sharks (Myers and Worm, 2003). Annual and decadal patterns of catches and their compositions can be assessed to understand the ability of the model to reproduce interannual and long-term changes in fisheries catches. Interpretation of fisheries catch data, however, must be done with care as changes or differences in fishing effort, gear, regulations, taxonomic identification, economics, or human behaviour can strongly affect the quantity, composition, and location of catches (Pinsky and Fogarty, 2012). For this reason, determining whether observed changes in catch data are caused by climate, ecology, or human behaviour can be complicated (e.g. the sablefish case study presented above). Fisheries-dependent data have substantial uncertainties because of inconsistent data quality and biases in sampling methods, timing, and location. Fisheries catches and landings data may be underreported (Zeller *et al.*, 2006), over-reported (Watson and Pauly, 2001) or misreported (Kleisner *et al.*, 2013), and the reliability and accuracy of the data may change over time. Also, biases in the location and timing of fishing activities render it challenging to standardize and use fisheries-dependent catch per unit effort data as an index of abundance (Maunder *et al.*, 2006). There may therefore be biases in using such data to interpret resource abundance and distribution (Walters, 2003).

Scientific survey data

Scientific surveys are useful across all three tiers of evaluation. They can provide spatial and temporal patterns of abundance, biomass,

biodiversity, and distribution. Among the benefits of scientific surveys is the use of standardized and repeatable methods, stratified random or fixed design to facilitate statistical inference, and documented survey locations so that both species presence and absence can be known. These properties make it more likely to attribute observed changes to particular drivers, such as fishing, pollution, and climate change, compared with fisheries data. For example, data from the California Cooperative Oceanic Fisheries Investigations (CalCOFI) (Bograd *et al.*, 2003) for the California Current Large Marine Ecosystem, which is strongly affected by decadal to multidecadal atmospheric oscillations, such as ENSO and PDO, provide detailed documentation of ecological changes since 1951. The CalCOFI data describe the abundance of plankton, including larval fish. A time series of larval fish abundance provide a useful proxy for adult fish abundance (Koslow *et al.*, 2013). Some surveys further record information on oceanographic conditions, which might be useful for simultaneously assessing the skill of the climatic and ecological components of LMR models. Although a number of surveys available have been sampling for more than four decades, care must be taken to ensure that large changes in survey methods have not biased the time series. A common standardization is to ensure that the same region has been surveyed consistently through time. Also, bias correction factors may be available to account for changes in survey methods (e.g. Ohman and Smith, 1995).

Although survey data can provide estimates of large-scale changes in the distribution of relative abundance or biomass of LMR (e.g. Pinsky *et al.*, 2013), they are regional in scale, typically conducted during a certain season, and are designed to sample a specific set of species or size classes (e.g. large groundfish). Different surveys also vary in time frame, and availability of long time-series survey data are limited. On the other hand, survey data are available for a range of ecosystem types (from the tropics to high latitudes), thereby allowing the examination of model performance across ecological gradients.

Species occurrence records

A major biological response to ocean changes is a shift in the distributions of marine species (Pinsky *et al.*, 2013; Poloczanska *et al.*, 2013), which can have further implications for marine ecosystems and LMR (Cheung *et al.*, 2010, 2013b). It is thus desirable for LMR models to realistically predict distributions for a wide range of species. A range of species distribution models have been applied to model LMRs under climate change (e.g. Jones and Cheung, 2015). The reliability of predicted species distributions

Table 3. Examples of observation datasets and the model metrics for comparison with LMR model outputs.

Data type/model metric	Time frame	Spatial aggregation	Taxonomic resolution	Examples of data sources
Fisheries data				
Total fisheries catch potential	Average from 1950–2010	Global and large marine ecosystems	Aggregated	SAU project
Species composition of catch (or the mean temperature of catch)	Annual from 1950–2010	Global and large marine ecosystems	Exploited taxa	SAU project
Catch of large pelagic fish	1970s–2000s	Global	Large pelagic fish (tunas)	Regional Fisheries Management Organizations for tunas and billfish, e.g. Myers and Worm (2003)
Survey data				
Rate of range shift of marine species	Average from 1970s to 2000s	Regional (North America continental shelf, North Sea)	By species of fish and invertebrates	Pinsky <i>et al.</i> (2013), ICES's International Bottom Trawl Survey
Community composition	1960s to 2010s	Regional (continental shelves around the world)	By species of fish and invertebrates vulnerable to bottom trawls	Worm <i>et al.</i> (2009)
Variability in abundance driven by large-scale oceanographic changes	1951–Present 1931–Present 1970s–2000s	Regional: California Current North Sea North America continental shelf	Larvae and eggs of exploited and unexploited fish Exploited and unexploited fish By species of fish and invertebrates	CalCOFI Moser <i>et al.</i> (2001) Continuous Plankton Recorder (CPR) Survey Pinsky <i>et al.</i> (2013)
Occurrence record	Mostly since the 20th century	Global	All marine taxa	OBIS Costello <i>et al.</i> (2007)

are often examined using geo-referenced species occurrence records and test statistics, such as the area under curve (AUC) of the receiver operating characteristics. These records are collated from a range of sources including museum collections, scientific expeditions and surveys, and fisheries records. Many are now publicly accessible through databases, such as the Global Biodiversity Information Facility (Robertson *et al.*, 2014) and the Ocean Biodiversity Information System (OBIS) (Costello *et al.*, 2007), and have frequently been standardized for taxonomy and checked for quality. Species occurrence records have the advantage in having a much broader spatial and taxonomic coverage than any single data source (e.g. from scientific survey only). However, problems with taxonomic misidentification, common names, synonyms, and errors in georeferencing are still present. Confidence in species occurrence data may also be reduced due to sampling bias (Webb *et al.*, 2010). Specifically, information on locations where unsuccessful sampling has occurred is not always available, making it difficult to determine the areas where specific species are absent and therefore to interpret test statistics such as the AUC (Pearce and Boyce, 2006).

To help inform the use of uncertain observational data in assessing model projections, a framework has been proposed to systematically assess the level of uncertainty associated with observational data particularly for climate change impact assessment (O'Connor *et al.*, 2015). This framework is based on evidence combined from theory, experiments and historical data with statistical analysis being undertaken to attribute any signals in observational data to climate change, thereby building confidence in the model. Such a framework will help identify cases where observational data are too uncertain to help assess model outputs, e.g. with insufficient temporal and spatial coverage of observational data to reveal underlying trends and patterns.

Post-processing of LMR model outputs is generally needed before they can be compared with empirical data, as there will inevitably be differences between LMR models due to variations in model structure and other factors. For example, output from species-based LMR models will be more directly comparable with empirical data. However, species-based LMR models may only include a subset of species or taxonomic groups that are included in the empirical data. In contrast, output from size-based models can easily be compared with aggregated LMR production. However, the lack of explicit representation of taxonomic identity in size-based models makes their output difficult to compare with species- or population-specific data. Approximations can be made in some cases to convert information from size- or trophic-based models into taxonomic-based data. For example, the abundance and production of organisms at size > 1 m can be assumed to represent adult large pelagic fish and can thus be compared with data from pelagic longline catches. Functional group-based LMR models are intermediate between species- and size-based models, and their outputs can be approximately converted to both taxonomic- or size-based aggregations. Thus, having identified the dominant taxonomic groups in a functional group, the dynamics of that functional group can be assumed to be representative of that taxonomic group. Functional groups that represent specific taxonomic groups of interest can also be included explicitly in the model (deYoung *et al.*, 2004; Griffith and Fulton, 2014).

From quantifying uncertainty to assessing risk

Given the large sources of uncertainty discussed in previous sections, a systematic exploration of potential future LMR states and the associated uncertainties is an important step towards a full

risk assessment that would allow us to understand the potential impact of climate change on human societies through, for example, diminished food security, income, or other ecosystem services. In general, risk consists of two components: (i) the magnitude of potential changes and (ii) the probability of occurrence of such changes. Previous climatic risk assessments have involved both quantitative risk-based approaches and more qualitative, social vulnerability approaches (Dessai and Hulme, 2004), or a combination of both (Brown *et al.*, 2012). Quantitative assessment generally involves identifying climate hazards and their probability of occurrence. For example, Li *et al.* (2009) assessed the drought risk for world crop production under climate change based on ensemble results from 20 general circulation model (GCM) and six emission scenarios. The ensemble of projections was used to estimate probability density functions of drought disaster frequency. Their results show a consistent increase in drought risk in the middle and end of the 21st century under climate change, leading to significant reductions in yield for major crops. In our case study of projecting changes in potential catches of sablefish in the Northeast Pacific (Figure 4), the probability of projecting a decrease in catch could be quantified by systematically exploring the envelope of uncertainty. Thorough estimates of risk can facilitate policy discussion for mitigation and/or adaptation in LMR management through the exploration of the potential for regrets/no-regrets policies and the associated costs and benefits (Polasky *et al.*, 2011). This approach to risk-based, ecosystem-based management has been developed for certain marine systems, for example in Australia (Hobday *et al.*, 2011). One area of risk assessment that remains particularly difficult to accurately quantify and yet important for guiding societal choices is an understanding of “tail risk”, or risk from extreme and high-impact, but low-probability, events (Weitzman, 2011).

Future direction of climate-LMR projections

The many sources of uncertainty in climate-LMR projections and computational cost will always limit our ability to fully explore uncertainty in climate-LMR projections. However, the framework described here provides a basis for concerted effort to improve estimation of uncertainty ranges for climate-LMR projections and, eventually, reduce these ranges. As was the case for physical climate projections, a climate-LMR ensemble offers a starting point. Systematic exploration of uncertainty space to identify prominent components for a given spatial scale, time horizon, and variable of interest can guide research investment and accelerate progress toward more accurate estimates of uncertainty bounds. More rigorous and standardized comparison with observations (i.e. model metrics) must also play a central role in building confidence in projections. In combination, these steps should produce more robust risk estimates for policy formulation that will promote LMR sustainability in a changing climate.

While adoption of the framework described herein will improve climate-LMR projections, many challenges must still be overcome. Various unknowns pose a major challenge to exploring the real scope of uncertainties. Particularly, adaptive responses in nature to climate change, and by society to changes in LMRs, are difficult to predict and are poorly understood (Pinsky and Fogarty, 2012). There are also “unknown-unknowns”, such as ecological tipping points, which contribute to uncertainties and that cannot be assessed with our current knowledge. This problem could be partly addressed by developing scenarios that aim to explore the sensitivity of outputs to such uncertainties, such as a scenario incorporating high levels of biological and social adaptation.

Additionally, when exploring structural uncertainty of the models, the sample of model structures is often assembled opportunistically based on existing models rather than strategically based on a systematic sampling of all plausible model structures. Furthermore, different climate-LMR models may not be entirely independent from each other as the models may be parameterized with similar datasets. This may result in biases in assessing the effects of model uncertainties on projections (Hawkins and Sutton, 2009). On the other hand, an ensemble of opportunities would be the most practical way to tackle the challenge of quantifying climate-LMR projection uncertainties and would help examine whether there is a need for large-scale cooperative initiatives that provide substantial resources and facilities to address these challenges.

Observational data that are available for comparison with LMR models generally only cover a short period and a limited number of regions. This magnifies the issues regarding uncertainties associated with observation errors, making it more challenging to attribute the reasons for any discrepancies between observations and model predictions. Moreover, many LMR models use available observational data for parameterization, thus the scope of using additional data for model testing is limited. Careful selection of statistical and cross-validation techniques can help mitigate this problem (Arlot and Celisse, 2010). Further discussion and consensus among LMR modellers is needed to develop criteria to identify unrealistic models (i.e. what type and how many discrepancies are needed before a model is excluded from an ensemble). These challenges reiterate the need to improve the sharing of observational data between scientists, institutes, and countries and develop data facilities to support their use in testing climate-LMR projections (Hollowed *et al.*, 2013).

Scenario development has not matured for LMR assessment. Scenarios specifically tailored for marine-related sectors are very limited, while existing assessments adopt scenarios that are used for more general-purposes (Millennium Ecosystem Assessment, 2005). These scenarios may not account for key uncertainties in the projected pathways of LMRs. In relation to this, fisheries models linking fishing to changes in LMRs, and the socio-economic conditions that are used to generate LMR scenarios are only starting to be developed for global- and basin-scale LMRs, although much effort has focused on regional- and local-scale fishing fleet dynamic models (van Putten *et al.*, 2012) and management strategy evaluation models. All existing global- or basin-scale LMR models either do not have explicit fisheries components or have simple assumptions of stock- or region-specific fishing mortality rates. Only recently has a global scale LMR model study included a spatially explicit fishing dynamics model to simulate changes in fishing effort (Christensen *et al.*, 2015). However, there is a need to improve efforts such as this to develop additional LMR-specific scenarios representing human activities before meaningful comparison of scenario uncertainties can be undertaken.

Understanding where uncertainty comes from and how it interacts with model components is necessary to improve the interpretation of model projections and to inform policy. Improving the quantification of uncertainties will therefore be a major area of development in climate-LMR projections to inform global and regional assessments of climate change impacts, vulnerability, and adaptation on marine ecosystems and related sectors.

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